

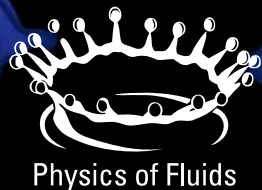
Convection-driven porous media flows:
Implications for carbon dioxide
sequestration

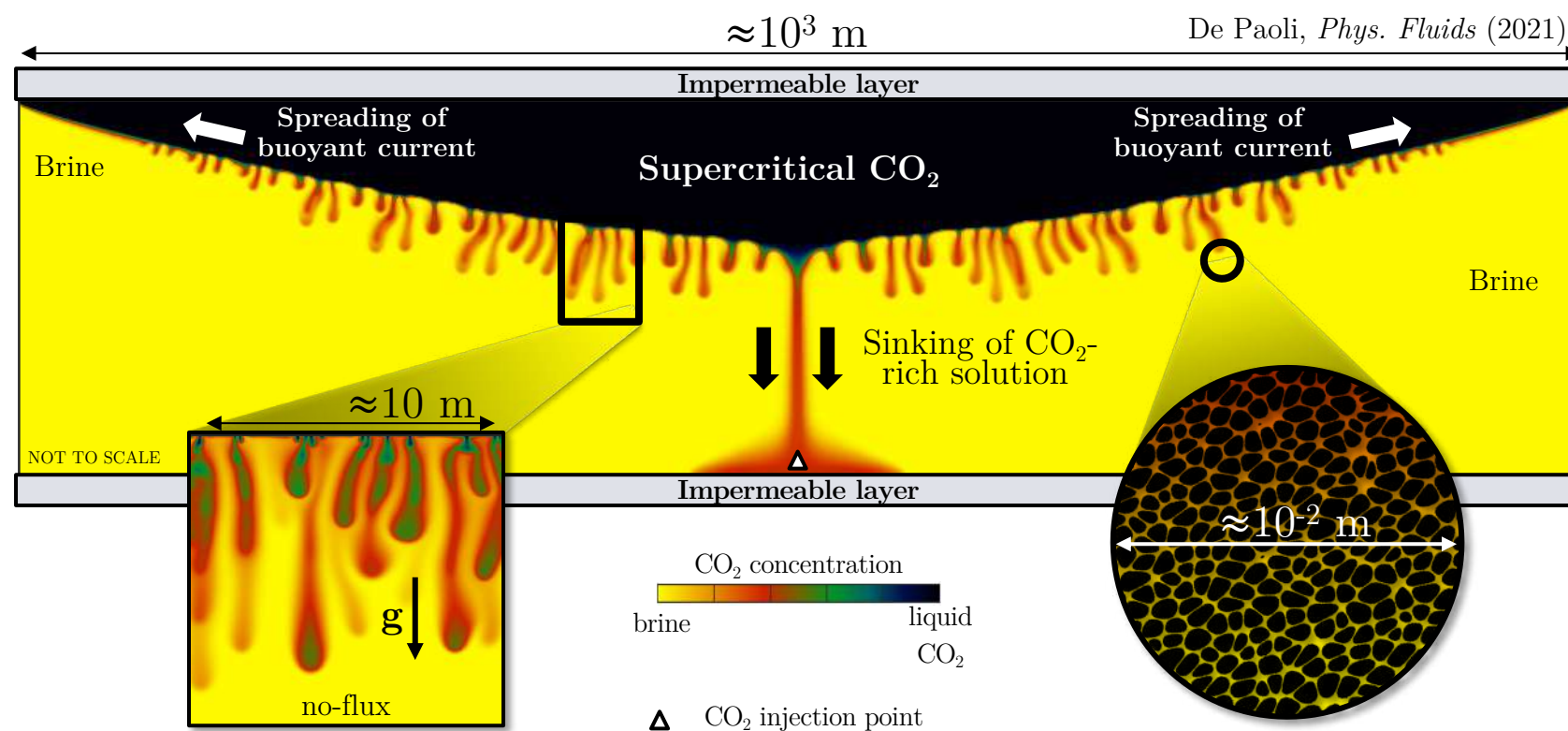
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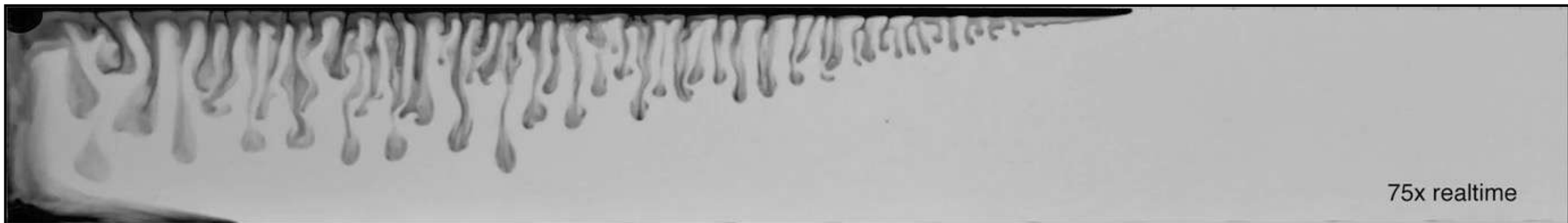


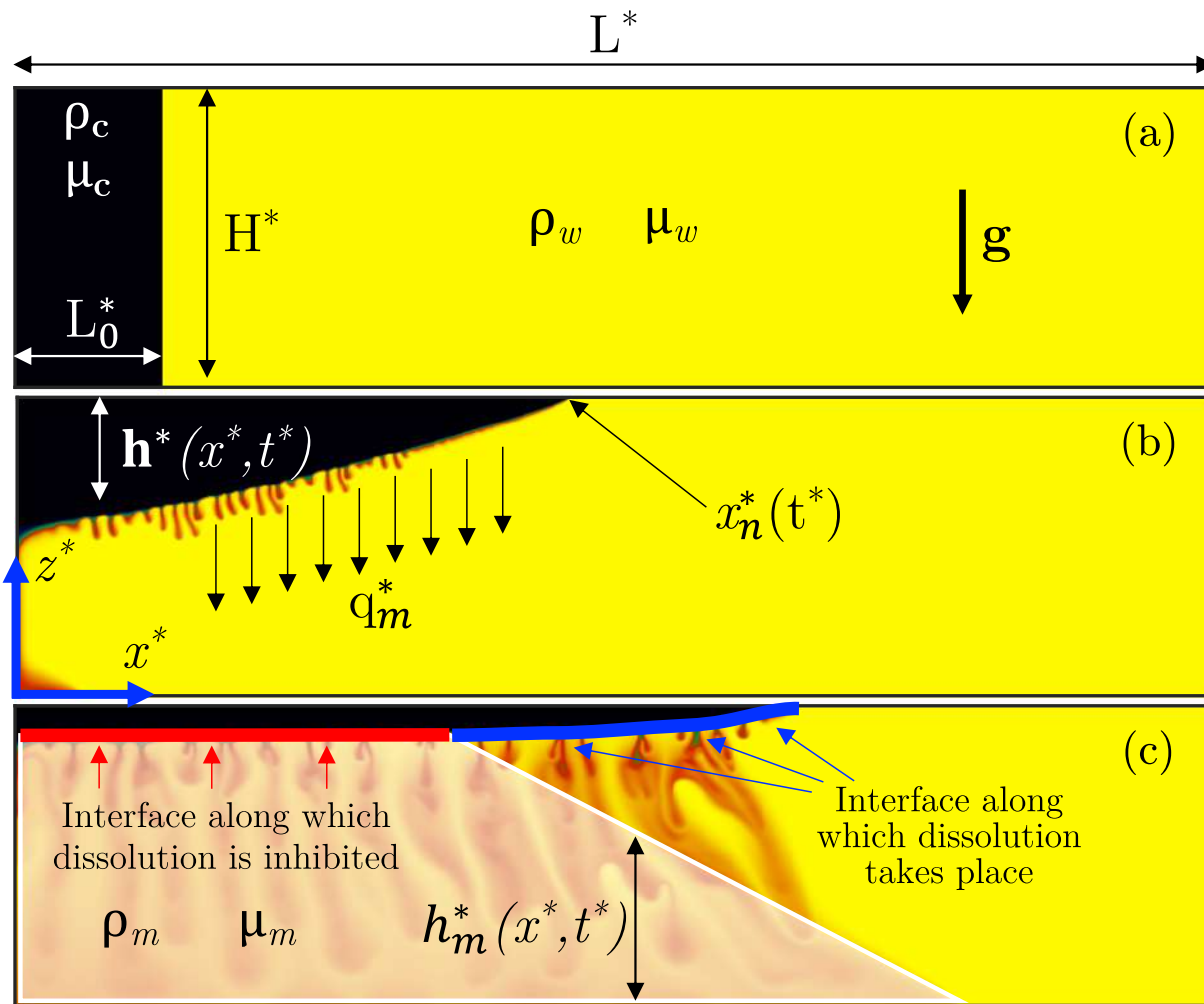


Reservoir properties

- anisotropy and heterogeneities
- finite size of confining layers
- effects of rock properties (mechanical dispersion)
- chemical dissolution and morphology variations
- ...

MacMinn et al., *Geophys. Res. Lett.* (2013)





De Paoli, *Phys. Fluids*. (2021)

$$\frac{\partial h}{\partial t} - \frac{\partial}{\partial x} \left[(1-f)h \frac{\partial h}{\partial x} - \delta f h_m \frac{\partial h_m}{\partial x} \right] = -\varepsilon_0,$$

$$\frac{\partial h_m}{\partial t} - \frac{\partial}{\partial x} \left[\delta(1-f_m)h_m \frac{\partial h_m}{\partial x} - f_m h \frac{\partial h}{\partial x} \right] = \frac{\varepsilon_0}{X_v}$$

$$f = \frac{Mh^*/H^*}{(M-1)h^*/H^* + (M_m-1)h_m^*/H^* + 1},$$

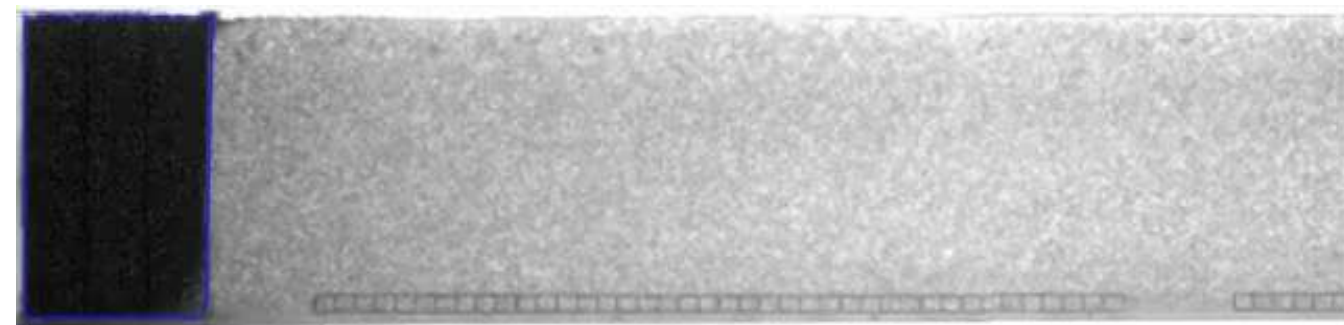
$$f_m = \frac{M_m h_m^*/H^*}{(M-1)h^*/H^* + (M_m-1)h_m^*/H^* + 1},$$

MacMinn, Neufeld, Hesse,
and Huppert, *Water Resour. Res.* (2012)

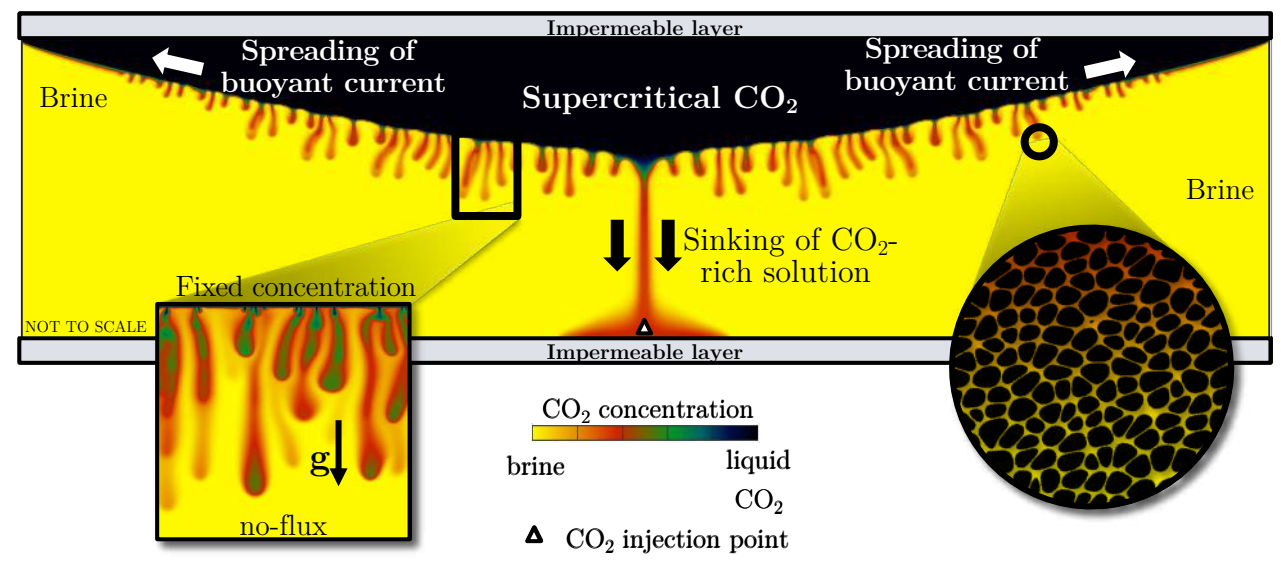
Mobility ratios $M = \mu_w/\mu_c$ and $M_m = \mu_w/\mu_m$

Buoyancy velocity ratio $\delta = W_m^*/W^*$

Volume fraction $X_v = \rho_m X_m / \rho_c$



MacMinn, Neufeld, Hesse, and Huppert, *Water Resour. Res.* (2012)



$$\frac{\partial h}{\partial t} - \frac{\partial}{\partial x} \left[(1 - f)h \frac{\partial h}{\partial x} - \delta f h_m \frac{\partial h_m}{\partial x} \right] = -\varepsilon_0$$

$$\frac{\partial h_m}{\partial t} - \frac{\partial}{\partial x} \left[\delta(1 - f_m)h_m \frac{\partial h_m}{\partial x} - f_m h \frac{\partial h}{\partial x} \right] = \frac{\varepsilon_0}{X_v}$$

$$f = \frac{Mh^*/H^*}{(M - 1)h^*/H^* + (M_m - 1)h_m^*/H^* + 1},$$

$$f_m = \frac{M_m h_m^*/H^*}{(M - 1)h^*/H^* + (M_m - 1)h_m^*/H^* + 1},$$

$$\varepsilon_0(x) = \begin{cases} 0 & \text{if } h(x) = 0 \text{ or } h(x) + h_m(x) = 1 \\ \varepsilon & \text{else,} \end{cases}$$

$$\varepsilon = \frac{q_m^*}{\phi W^*} \left(\frac{L_0^*}{H^*} \right)^2$$

How to determine the dissolution rate q_m^* ?

Dimensionless equations

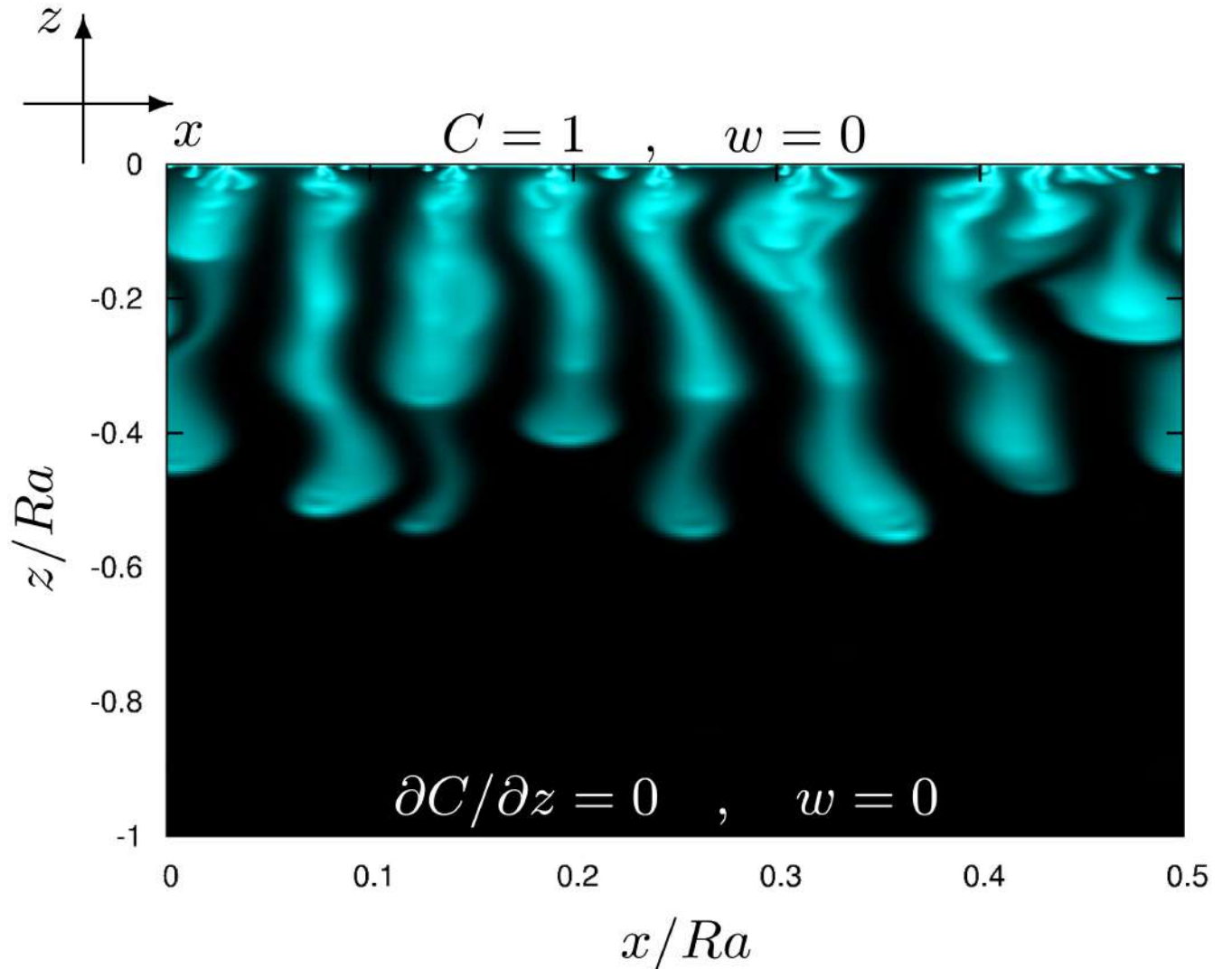
$$\frac{\partial C}{\partial t} + u \frac{\partial C}{\partial x} + w \frac{\partial C}{\partial z} = \frac{1}{Ra} \left(\frac{\partial^2 C}{\partial x^2} + \frac{\partial^2 C}{\partial z^2} \right)$$

$$u = -\frac{\partial P}{\partial x}, \quad w = -\frac{\partial P}{\partial z} - C$$

$$\frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0$$

Governing parameter

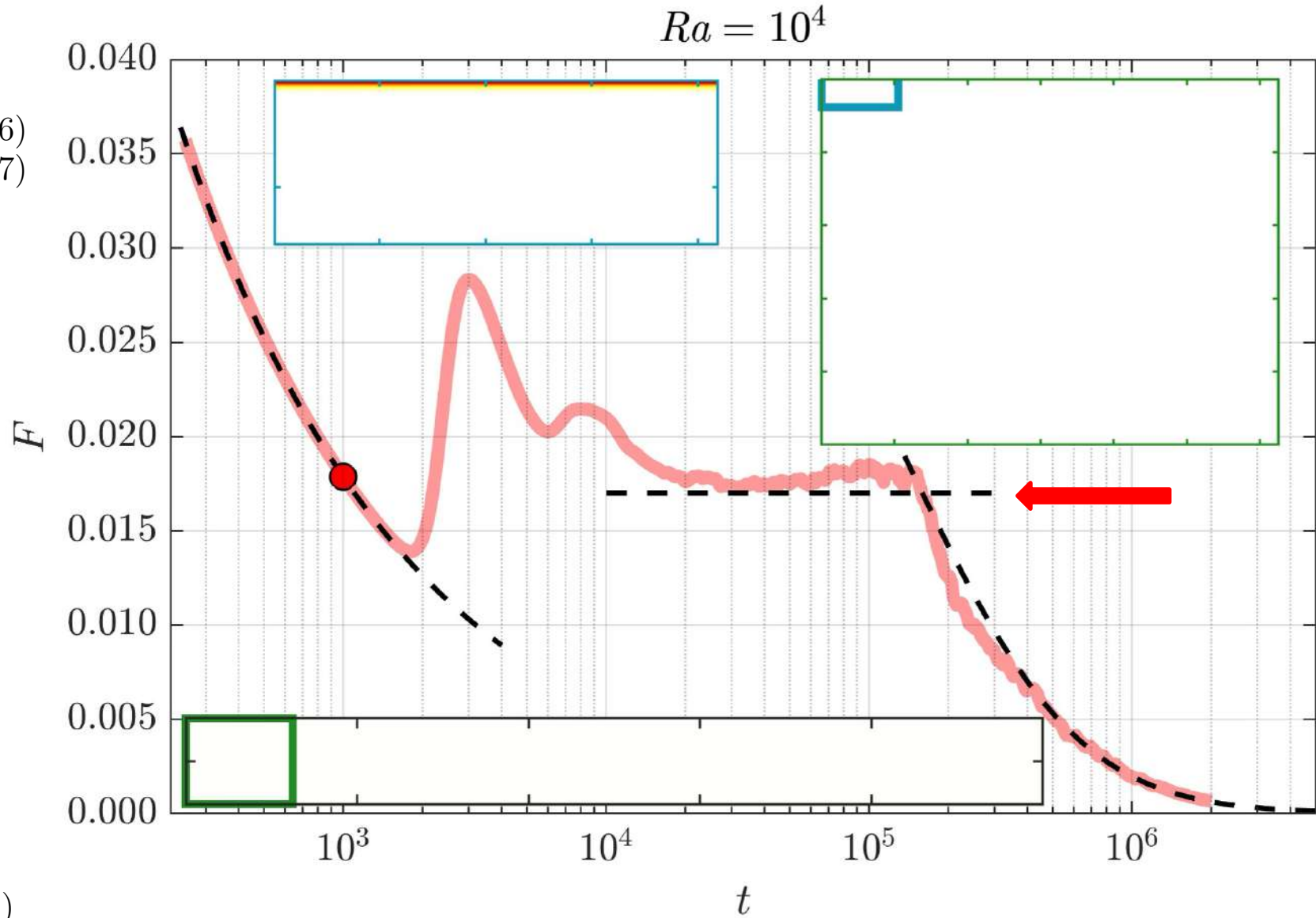
$$Ra = \frac{gH^*k_v\Delta\rho^*}{\mu\Phi D}$$



De Paoli, Zonta and Soldati, *Phys. Fluids* (2016)
 De Paoli, Zonta and Soldati, *Phys. Fluids* (2017)

$$F(t) = \frac{1}{L} \int_0^L \left. \frac{\partial C}{\partial z} \right|_{z=0} dx$$

Examples of model extension:
 effect of **anisotropy** of the medium



See also Slim, *J. Fluid Mech.* (2014)
 Hewitt, Neufeld & Lister, *J. Fluid Mech.* (2013)

Convection in anisotropic media

Examples of model extension:
effect of **anisotropy** of the medium

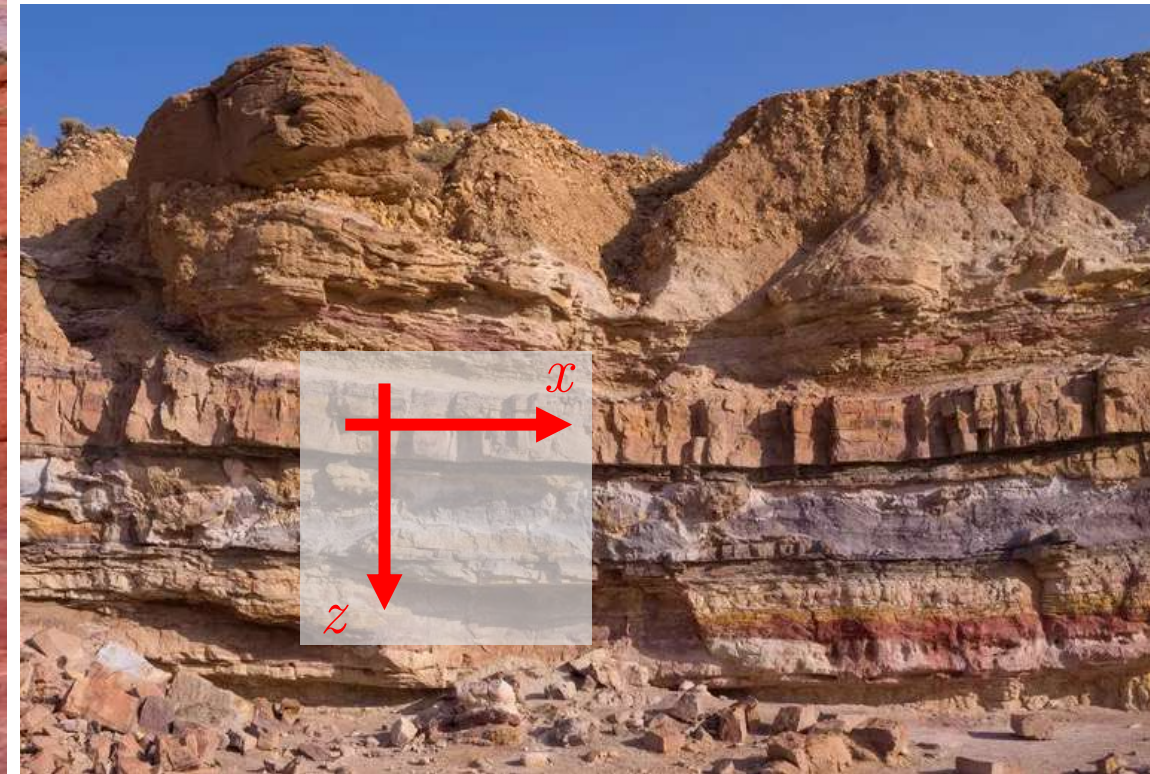
Sedimentary rocks: Rocks formed by stratification

Assumptions:

1. Homogeneous porous medium
2. Anisotropic porous medium
 - Principal directions of the permeability tensor aligned with the reference frame

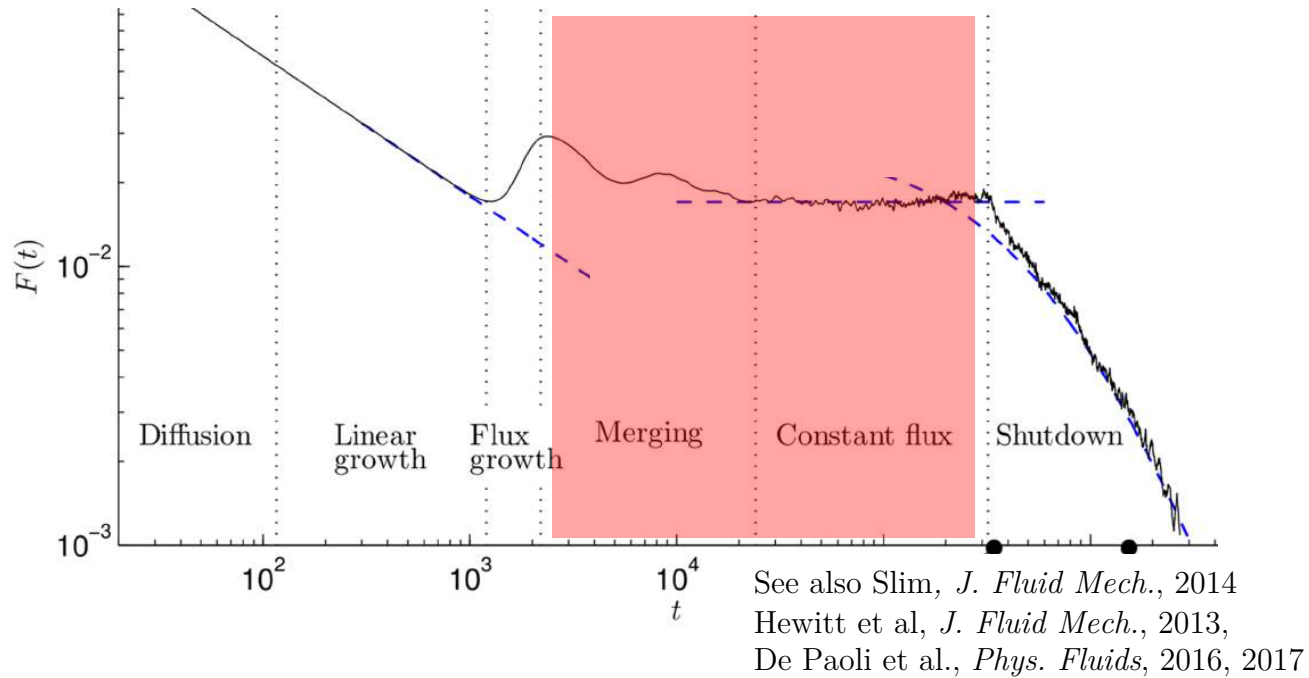


benedek / Getty Images

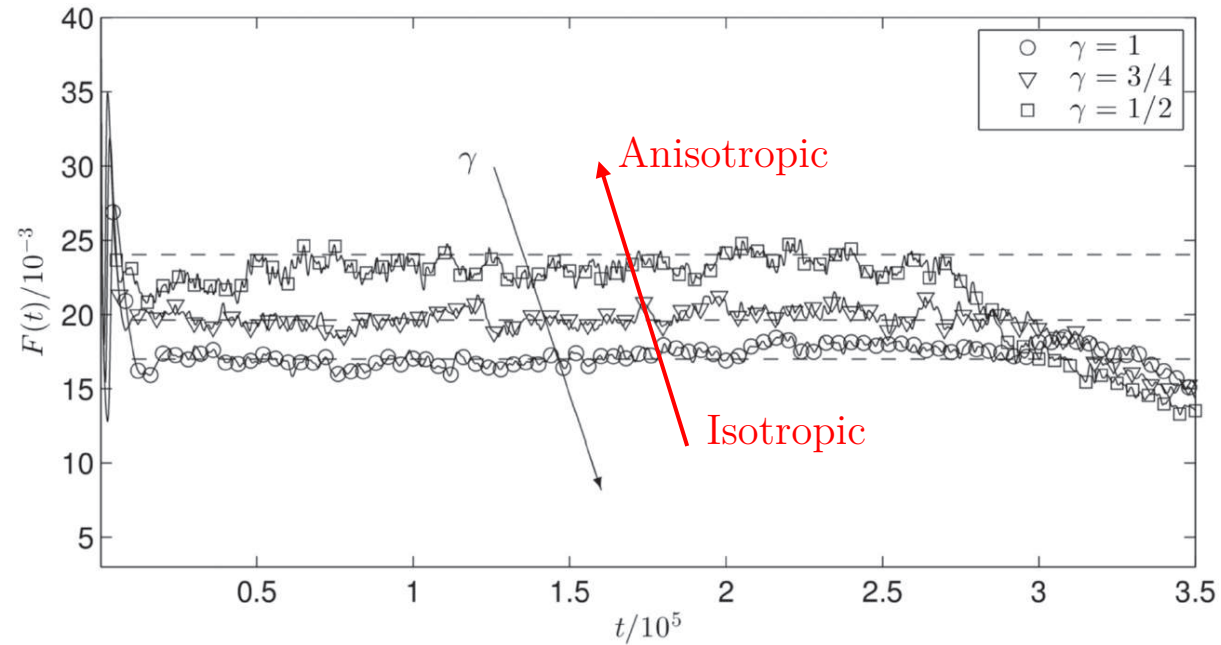


Rhododendrites/Wikimedia Commons/CC BY 4.0

isotropic medium



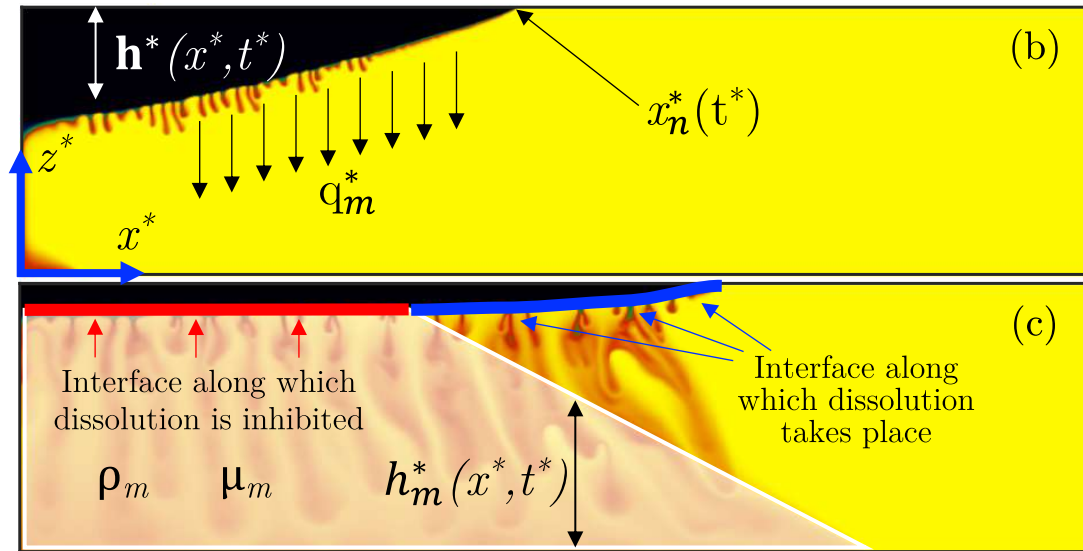
$$F(t) = \frac{1}{L} \int_0^L \frac{\partial C}{\partial z} \Big|_{z=0} dx$$



Convection-dominated $F(t) = 0.017$

Strong influence of γ on flux

$$q_m^* \equiv F(t) = 0.017\gamma^{-1/2}$$



Darcy-scale simulations:



$$\text{dissolution rate } q_m^* \sim \gamma^{-\frac{1}{2}}$$



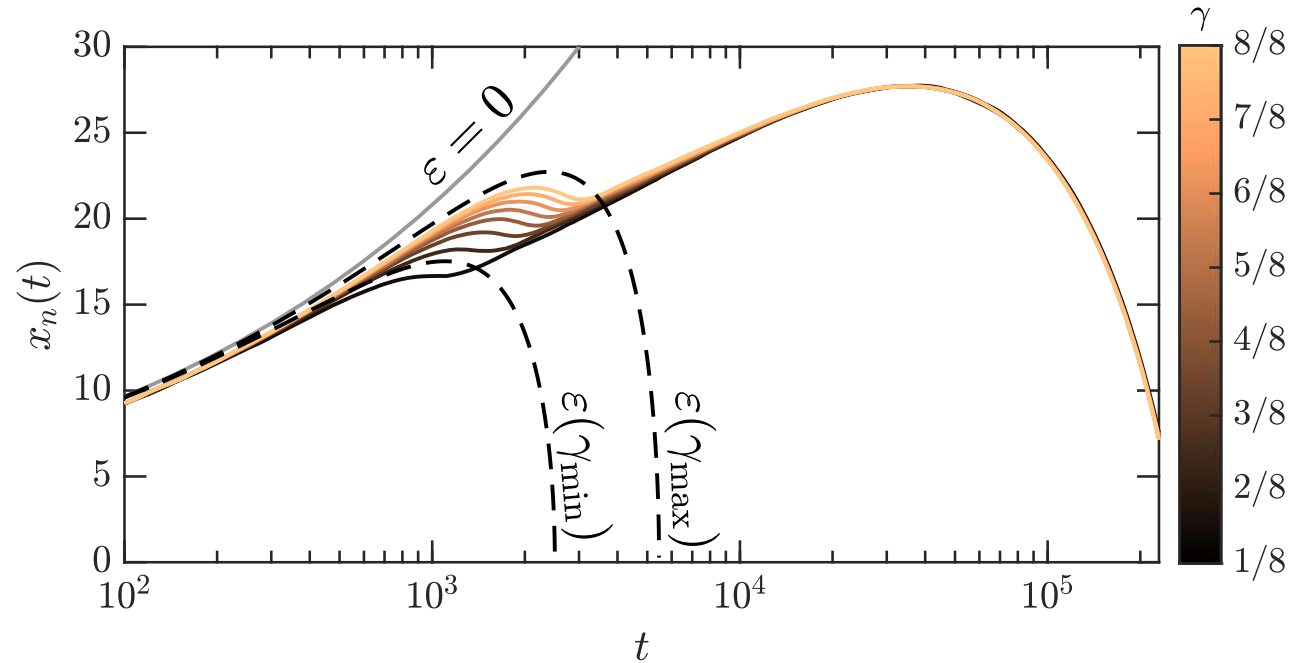
dissolution increases with the anisotropy of the medium

Sedimentary rocks are anisotropic

$$\gamma = \frac{k_v}{k_h} < 1$$

$\gamma = 1$ isotropic

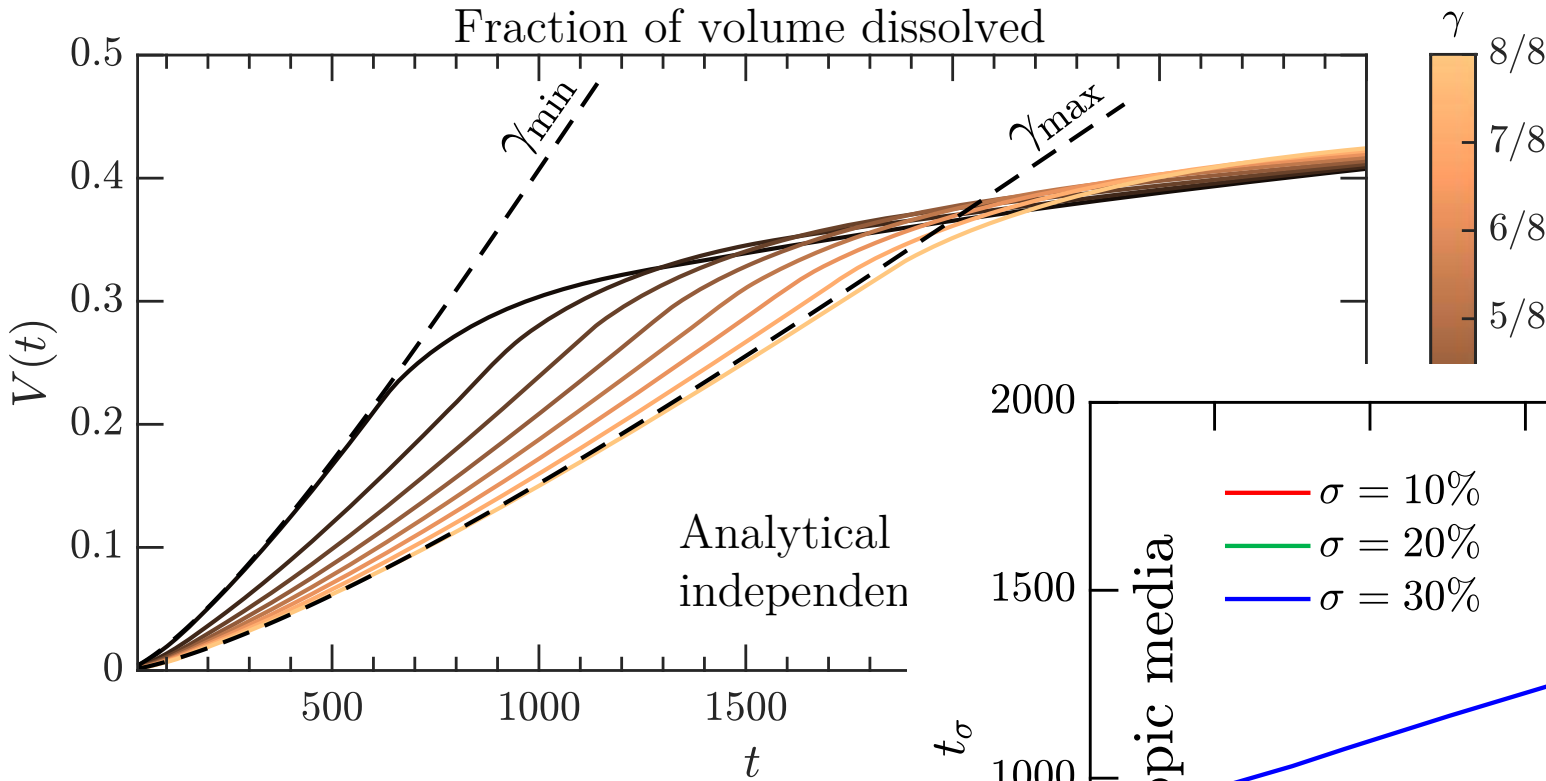
$\gamma = 1/8$ strongly anisotropic



Analytical solution in case of

- no-dissolution ———
- independent currents - - - - -

Effect of anisotropy

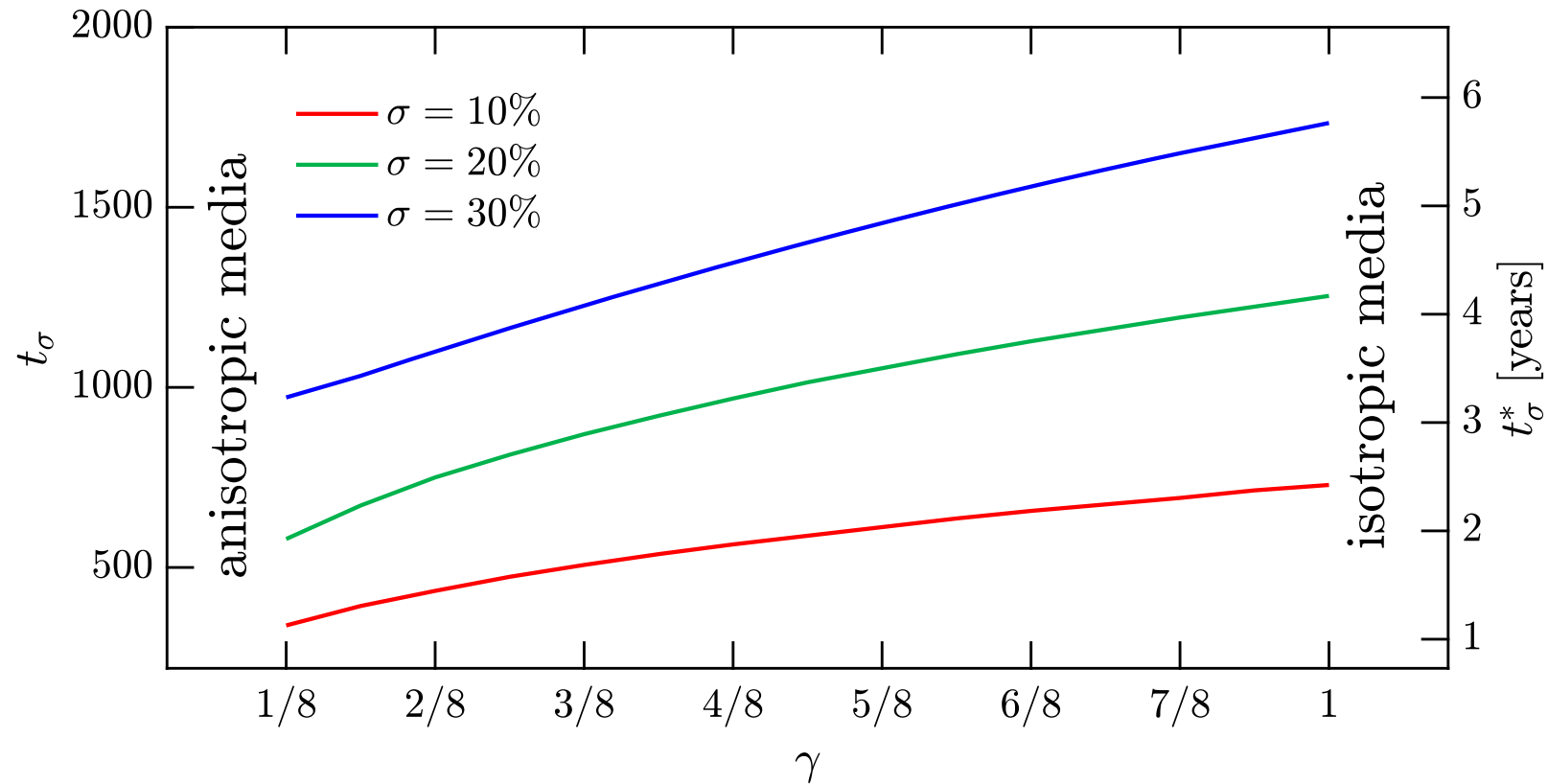


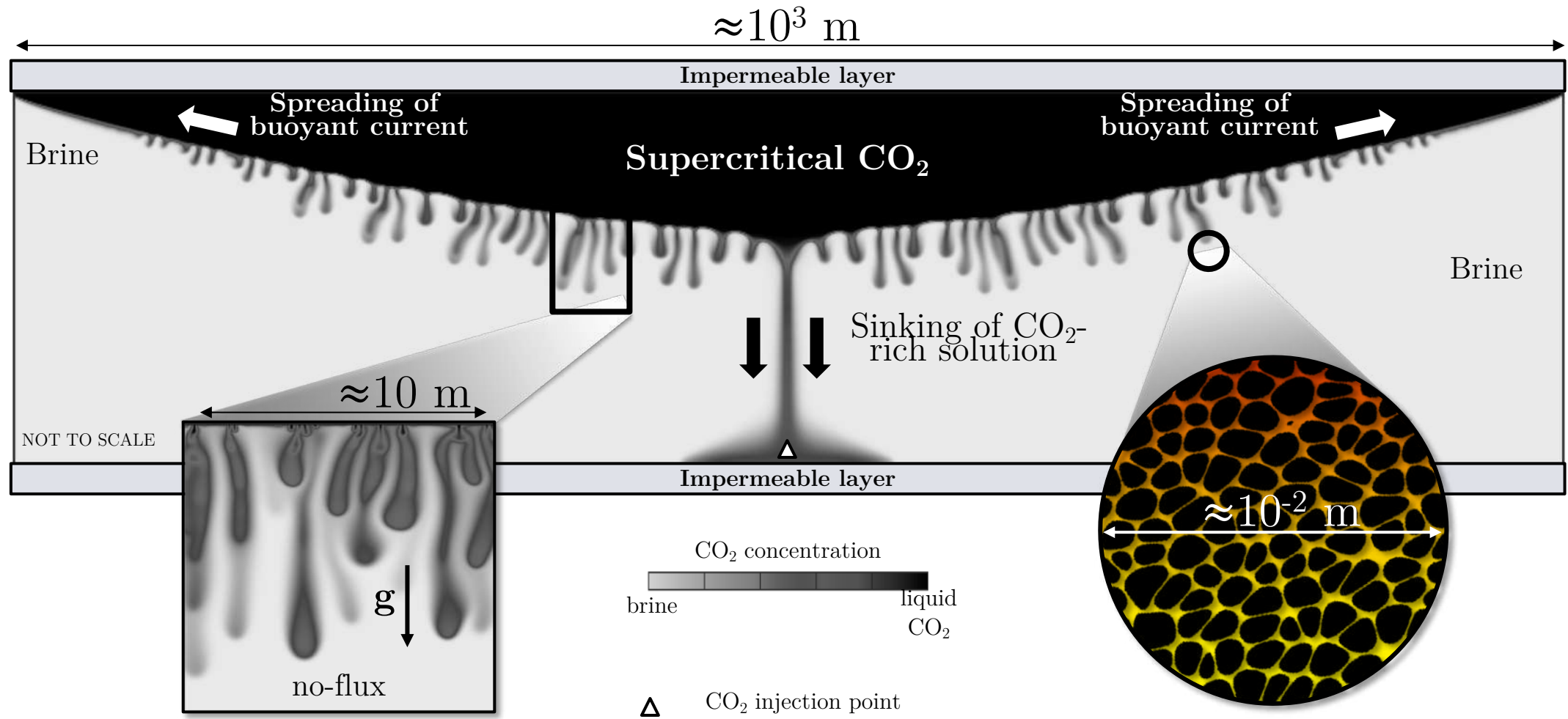
Sedimentary rocks are anisotropic $\gamma = \frac{k_v}{k_h} < 1$

$\gamma = 1$ isotropic

$\gamma = 1/8$ strongly anisotropic

Long-term predictions
of real-scale reservoirs





This research was funded in part by the Austrian Science Fund (FWF) [Grant J-4612]

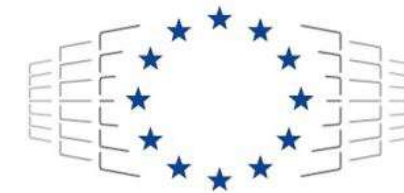


Der Wissenschaftsfonds.

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MAX-PLANCK-GESELLSCHAFT



EuroHPC
Joint Undertaking

High-resolution images, movies and slides are available upon request to m.depaoli@utwente.nl